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INTRUSION OF RECENT AIR IN POLAR STRATOSPHERE DURING SUMMER 2009 REVEALED BY BALLOON-BORNE IN SITU CO MEASUREMENTS

Gisèle Krysztofiak⁽¹⁾, Rémi Thiéblemont⁽¹⁾, Valéry Catoire⁽¹⁾, Nathalie Huret⁽¹⁾, Sébastien Payan⁽²⁾, Fabrice Jegou⁽¹⁾, Marc Antoine Drouin⁽¹⁾, Claude Robert⁽¹⁾.

¹ LPC2E-OSUC, UMR 6115 CNRS-Université d'Orléans, 3A Avenue de la Recherche Scientifique, 45071 Orléans Cedex 2, France, Email: gisele.krysztofiak@cnrs-orleans.fr

² LPMAA, UMR 7092 CNRS-Université Pierre et Marie Curie, 3 rue Galilée 94200 Ivry sur Seine, France.

ABSTRACT

The SPIRALE (Spectroscopie Infa-Rouge par Absorption de Lasers Embarqués) balloon-borne instrument has been launched twice within 17 days in the polar region (Kiruna, Sweden, 67.9°N - 21.1°E) during summer, at the beginning and at the end of August 2009. In situ measurements of several trace gases have been performed including CO and O₃ between 10 and 34 km height, with very high vertical resolution (~5 m). The both flight results are compared and the CO stratospheric profile of the first flight presents specific structures associated with mid-latitude intrusion in the lowest stratospheric levels. Their interpretation is made with the help of results from several modeling tools (MIMOSA and FLEXTRA) and available satellite data (IASI). We also used the O₃ profile correlated with CO to calculate the proportion of recent air in the polar stratosphere. The results indicate the impact of East Asia urban pollution on the chemistry of polar stratosphere in summer.

1. INTRODUCTION

Arctic atmosphere has been believed to be very clean for long time. Nowadays the polar atmosphere become more and more polluted due to the increase of the pollutant emissions and the long scale transport [1, 2]. Carbon monoxide (CO) is one of these pollutants with a long chemical lifetime (~2 months in the troposphere), allowing for a long scale study. Industrial processes, incomplete combustion, biomass burning and methane oxidation are the principal sources of CO in the troposphere and reaction with OH is the main sink. CO plays an important role in the tropospheric oxidation capacity and in the production of tropospheric ozone, depending on the NO_x quantities. Depending on the latitudes, the typical tropospheric CO mixing ratio can vary from 40 to 200 ppbv [3] with higher values in the northern hemisphere than in the southern hemisphere. Most of the previous studies on the transport of CO across the Pacific Ocean and into the polar

atmosphere used satellite measurements or airborne data [4, 5]. In situ measurements from SPIRALE balloon-borne instrument were used in this study on 7 August and 24 August 2009 near Kiruna (67.9°N-21.1°E) in Sweden. SPIRALE flights were part of the balloon campaign conducted by CNES within the frame of the StraPolÉté project funded by the French agencies ANR, CNES and IPEV, contributing to the International Polar Year. In this paper, the method used to perform this study is summarized in Section 2, Section 3 shows the CO observations made during the two SPIRALE flights, and their intercomparisons with model and satellite data.

2. DATA AND METHODOLOGY

2.1 SPIRALE instrument

SPIRALE (Spectroscopie InfraRouge par Absorption de Lasers Embarqués) is a tunable diode laser spectrometer for in situ measurements of trace gas species from the upper troposphere to the middle stratosphere (~34 km height). A detailed description of the instrument can be found in a previous paper [6]. In brief, the absorption of six lasers takes place between two mirrors distant of 3.5 m in a multipass optical Herriott cell located at the extremities of a deployable mast below the gondola, leading to 434 m of pathlength. SPIRALE has a rapid measurement acquisition (1.1 s), a very good accuracy (3% to 30%, depending on the abundance of the species) and a low detection limit (sub-ppbv levels).

The SPIRALE balloon-borne instrument has flown on 7 August and during the night of 24 to 25 August 2009. During both flights, several species, such as O₃, N₂O, CH₄, HCl, NO₂, HNO₃ and CO were measured from 10-16 km (depending on the chemical abundances) to 34 km height.

2.2 IASI and TERRA/MODIS satellites

In order to detect the sources of carbon monoxide, we used measurements from IASI (Interféromètre Atmosphérique de Sondage Infrarouge) and

MODIS (Moderate Resolution Imaging Spectroradiometer) instruments onboard METOP and TERRA/AQUA satellites, respectively.

The polar-orbiting MetOp-A, launched on 19 October 2006, is the first of three successive MetOp satellites. The IASI instrument is a nadir-looking high resolution Fourier Transform Spectrometer (FTS). IASI is designed to provide atmospheric temperature and water vapour profiles for operational meteorology. Atmospheric concentrations for several key species important to climate forcing and atmospheric chemistry monitoring can be derived from IASI radiance measurements [7]. In addition, IASI offers an excellent horizontal coverage due to its across track swath width of 2200 km, allowing global coverage twice a day, with a field of view sampled by 2×2 circular pixels each with a 12 km footprint diameter. IASI measures carbon monoxide (CO) on a global scale, twice a day. CO total columns and vertical profiles are retrieved in near real time from the nadir radiance spectra measured by the instrument in the thermal infrared (TIR) spectral range [8].

MODIS [9] detects a wide spectral range of electromagnetic energy at three spatial resolutions every day. We use the fire and thermal anomalies from climate modeling grid fire products at 0.5° resolution. Fire detection is performed using a contextual algorithm [9] that exploits the strong emission of mid-infrared radiation from fires.

2.3 Models

Three models, namely FLEXTRA, REPROBUS and MIMOSA, have been used to calculate air masses trajectories, stratospheric chemistry, and potential vorticity advection, respectively. FLEXTRA [10] is a Lagrangian atmospheric trajectory model, developed at the Institute of Meteorology and Geophysics University of Vienna to compute trajectories from meteorological fields of the European Centre of Medium-Range Weather Forecasts (ECMWF). We used the ERA-interim model fields with horizontal resolutions of $1^\circ \times 1^\circ$, 60 vertical levels, and with a frequency of 3 hours. Clusters of three-dimensional backward trajectories (which means the three wind components are used) uniformly distributed in a volume of size $0.5^\circ \times 0.5^\circ$ along latitude and longitude and 500 m height are used.

Potential vorticity (PV) maps are calculated using the “Modèle Isentropique de transport Mésos-échelle de l’Ozone Stratosphérique par Advection” (MIMOSA) contour advection model [11]. This model performs high resolution advection calculations based on the ERA-Interim reanalysis [12] of wind, pressure and temperature fields. MIMOSA initially computes the PV field at a resolution of 1.125° in latitude and longitude (T106) vertically interpolated on an isentropic

surface. This field is then interpolated on an x-y grid centered on the North Pole with a horizontal resolution of 37×37 km (three grid points / degree) and advected with a time step of one hour. To preserve the homogeneity of the field, a regridding of the PV field on the original grid is done every 6 hours. The information on diabatic changes in the PV field at large scales can be extracted from the ERA-Interim fields. In MIMOSA model this is done by applying to the advected field a relaxation towards the ERA-Interim PV field calculations with a time constant of 10 days. This technique allows MIMOSA to run continuously over periods of several months in order to follow the evolution of dynamical barriers and fine scale structures such as vortex remnants and tropical intrusions [13, 14, 15].

REPROBUS 3D chemistry-transport model [16, 17] contains a detailed description of O_x , NO_x , HO_x , ClO_x , BrO_x and CHO_x chemistry. It calculates the chemical evolution of 55 species using 160 gas-phase reactions and 6 heterogeneous reactions. Reaction rates coefficients are taken from the recommendations of Sander et al. [18]. The photolysis rates are calculated at every time step using a look-up table from the Tropospheric and Ultraviolet visible (TUV) model [19]. REPROBUS model extends from the ground up to 10 hPa, with a vertical resolution varying from less than 1 km near the tropopause level to 2.2 km in the upper part of the stratosphere. The horizontal resolution used for this study is 2° latitude \times 2° longitude. Zonally symmetric initial tracer fields were taken from the two-dimensional model described by Bekki et al. [20]. Winds and temperatures (ERA-INTERIM) by the ECMWF were used during the simulation to drive the transport of the stratospheric species and to compute their loss and production rates, respectively.

3. POLAR STRATOSPHERIC CO IN AUGUST 2009 BELOW 380 K ISENTROPIC SURFACE

3.1 Observations

The vertical profiles of carbon monoxide (CO) derived from SPIRALE measurements for both flights are displayed in Fig. 1. The cold point tropopauses were located at 11.78 and 11.34 km height on 7 and 24 August, respectively.

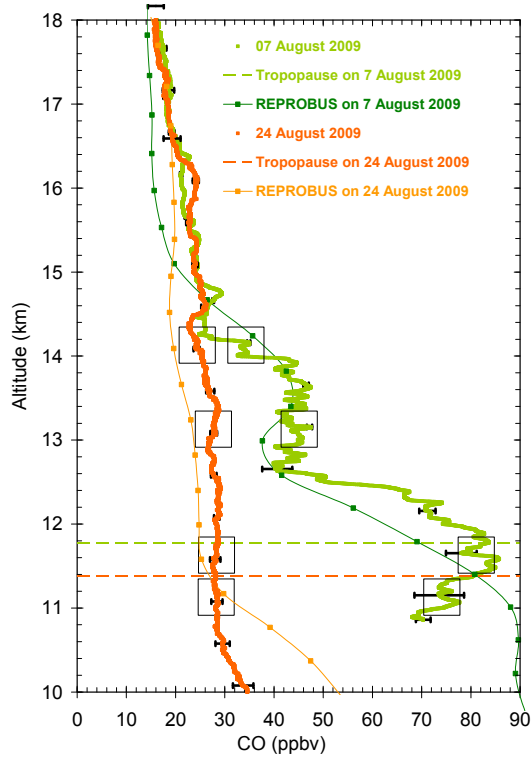


Figure 1. SPIRALE measurements of CO mixing ratios (in ppbv) with error bars over Kiruna (Sweden) on 7 August (green) and 24 August 2009 (orange) coupled with REPROBUS simulation. The dashed vertical lines represent the tropopause height for the two flights. The squares represent the altitudes of the clusters for the backward trajectories.

Although both flights were performed in August 2009 near Kiruna at 17 days interval, the vertical profiles of CO show a very strong difference below 14 km. CO mixing ratios are larger by ~50 ppbv at 11.6 km on 7 August than on 24 August 2009. On 7 August, the profile is mainly characterized by two layers, between 12.6 and 14 km with ~50 ppbv CO and below 12.4 km with ~80 ppbv CO. These ranges of concentrations are characteristic of mid-latitude concentration of CO. On the contrary, the vertical profile of the 24 August appears undisturbed with low values of CO (from 30 ppbv near the tropopause to 20 ppbv near 18 km) typical from polar latitudes.

The chemistry-transport model REPROBUS simulates also the different between the two flights. However the comparisons between each vertical profile and the associated simulation reveal strong discrepancies. The layer between 11 and 12.4 km (7 August) is not correctly resolved by the model which underestimates the value of CO. The underestimation of CO is also seen in general on 24 August profile and on the top of the profile on 7 August. Finally, below 11.2 km, the simulation overestimates the CO by 20 ppbv approximately.

This later bias could be improved by correcting the CO emission rate in the model.

The transport model MIMOSA reveals the presence of recent air from mid-latitude nearby Kiruna on 7 August 2009. Fig. 2 shows the intrusion of mid-latitude air, depicted by a low PV values tongue, toward polar latitudes at 340 K and 380 K isentropic surface levels, i.e. ~12 and ~14 km altitudes, respectively. In contrast, the 24 August PV maps point out typical polar air. The intrusion of lower latitude air masses on 7 August partly explains the high anomaly of CO in the vertical profile between 11 and 14 km. This is the reason why the origin of the air masses has to be investigated.

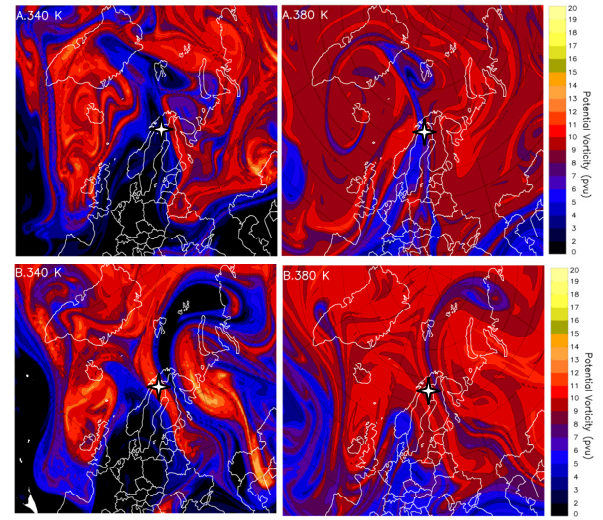


Figure 2. Potential vorticity from MIMOSA on the 340 K and the 380 K levels on 7 August 2009 at 00 UTC (A) and on 25 August 2009 at 00 UTC (B).

The white cross shows the location of the flights.

3.2 Origin of air masses

The FLEXTRA model allows us to know the origin of air masses. Ten day 3D backward trajectories are performed (shown in Fig. 3 below) for both flights of SPIRALE for the four clusters shown in Fig. 1. The backward trajectories reveal differences in the origin of air masses between both flights. For the flight on 25 August 2009, air comes exclusively from polar latitudes. For the flight on 7 August 2009, air masses originate from mid-latitudes and in particular from West Pacific ocean and East Asia. Fig. 3 also shows the altitude of the trajectories (using a color code) and reveals that, for the first flight, most of them come from low altitudes (< 6 km), which consequently suggests the presence of a convective area over the Pacific.

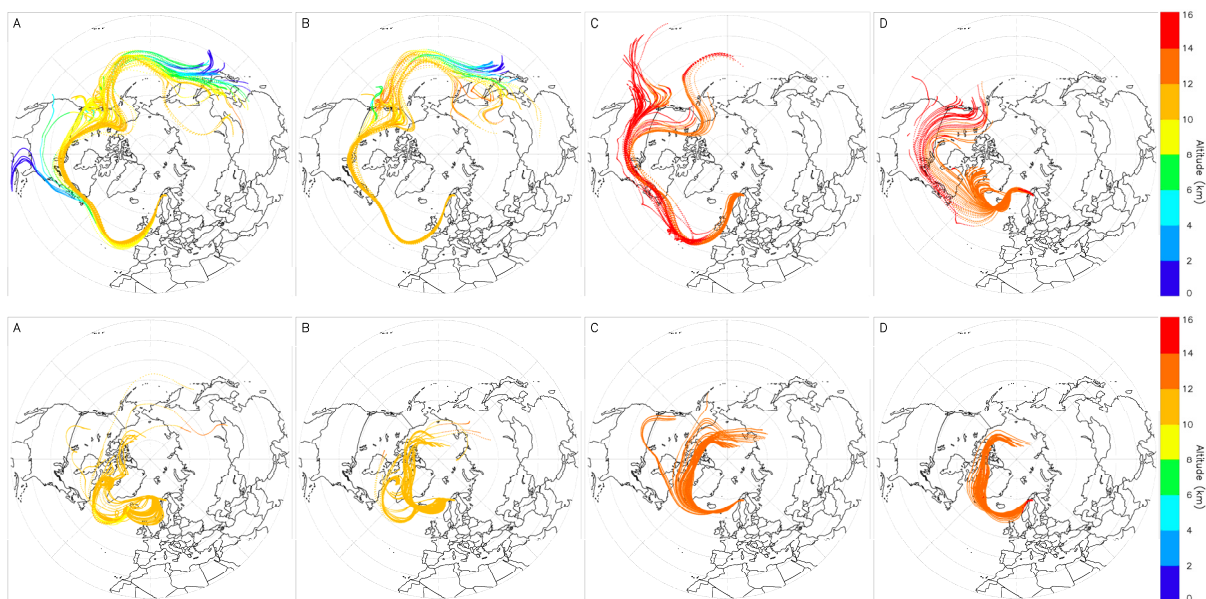


Figure 3. Ten day backward trajectories on 7 August 2009 (top) and 25 August 2009 (bottom) for clusters centred at A: 11130 m, B: 11630 m, C: 13130 m and D: 14130 m.

Further analyses of transport are discussed in Section 3.4. We are first investigating the origins of the pollution in the region.

3.3 Anthropogenic and natural sources of CO in Asia:

According to the backward trajectories, the air masses originate from the North East of Asia. So, we investigated the impact of emissions from boreal fires in North-East Russia and the anthropogenic sources from North-East China in the polar air in Kiruna on 7 August 2009.

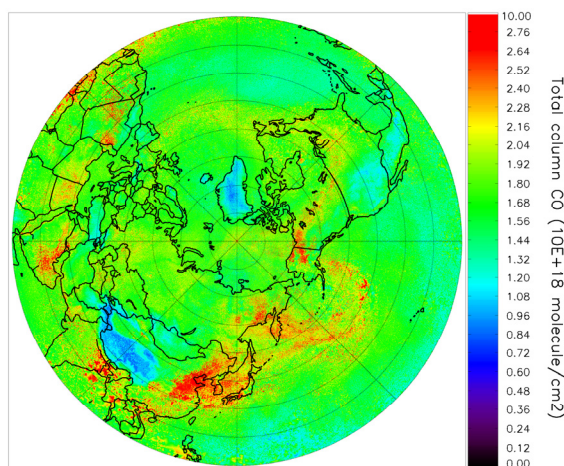


Figure 4. Carbon monoxide total column from IASI satellite from 27 July 2009 to 2 August 2009. Map is a polar stereographic projection with the Greenwich meridian at left.

IASI satellite measures the total column of CO (shown in Fig. 4 in molecule cm^{-2} from 27 July 2009 to 2 August 2009 (10 to 5 days before the

flight). IASI reveals two areas of strong concentration of CO over North East China and over North East of Russia. Moreover, according to Jiang et al. [21] the CO concentration in the upper troposphere is more important in summer due to strong convection and surface emission. Jin et al. [5] indicated the maximum of CO is located between the equator and 40°N in boreal spring and between 30°N and 50°N in July - August. The migration of the CO belt from spring to summer is due to the northward movement of the subtropical westerly jet. Unfortunately, IASI does not allow to identify the origin of the pollution (anthropogenic and/or fires). So, we decided to use the results of MODIS instrument on the AQUA and TERRA satellites to look for fires as possible causes of CO high levels in the lower polar stratosphere. Fig. 5 shows the detected fires. MODIS detected very strong fires in North East of Russia.

So, two areas with two different pollutions are detected, the first over North-East China and Korea with mostly anthropogenic sources from megalopolises and the second over North-East Russia with strong forest fires sources. According to Jin et al. [5] Asian fossil fuel is the biggest source of upper tropospheric CO over the north Pacific up to 27-30% in the NH summer and biomass burning contribute only for less than 10% to the total CO in the upper troposphere. These source regions of CO pollution are qualitatively consistent with the origin of the backward trajectories for the 7 August 2009 with most of the backward trajectories coming from the megalopolises.

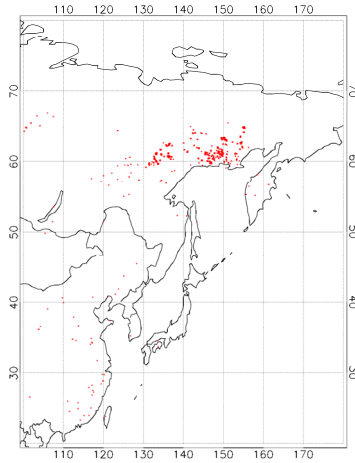


Figure 5. Fire detection by MODIS over East Asia from 27 July 2009 to 2 August 2009 (each red point symbolizes a fire location).

3.4 Transport of mid-latitude air into the polar stratosphere

Fig. 6 shows the horizontal wind at 200 hPa and the geopotential at 500 hPa on 29 July 2009 coupled with the ten day backward trajectories on 7 August 2009. The strong horizontal wind reveals the presence of the jet stream and the geopotential height reveals the presence of depression and anticyclone areas. The altitude of the backward trajectories reveals that some of the air masses undergo fast vertical transport from western Pacific before entering in the atmospheric circulation via the jet stream. Moreover, the long transport of pollutants from Asia across the Pacific is stronger in boreal spring and summer [5, 21]. Near Europe, the air masses undergo an isentropic transport across the subtropical barrier.

To summarize, after emission at ground by anthropogenic sources and/or fires, air masses containing CO undergo fast vertical transport to the upper troposphere, then cross the Pacific Ocean and the North America along the jet stream and finish by entering the polar stratosphere by isentropic transport.

It is well known [22] that this region, where isentropic surfaces intersect the tropopause, is favorable to the horizontal mixing of stratospheric and tropospheric air. Tracer-tracer correlations are a powerful diagnosis for stratospheric transport and mixing. In particular, we are interested in the O_3 -CO correlation in the upper troposphere - lower stratosphere (UTLS) that has also been used in recent studies for transport diagnoses [23, 24].

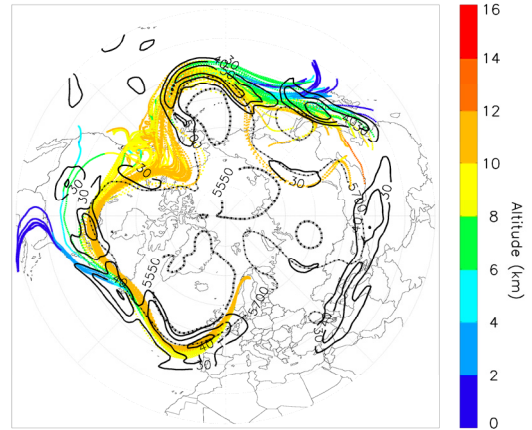


Figure 6. Ten day backward trajectories for a cluster centred at 11130 m on 7 August 2009 coupled with the ERA-Interim horizontal wind module (m/s) (solid line) at 200 hPa and geopotential height (m) (dashed-dotted line) at 500 hPa on 29 July 2009.

Fig. 7 shows an example of O_3 -CO correlations for two different latitudes. The first one is obtained from a previous balloon flight in June 2005 with the same instrument in a non polar region (Brazil, 5.1°S-42.9°W) and the second one is obtained from the polar flight measurements studied in the present paper.

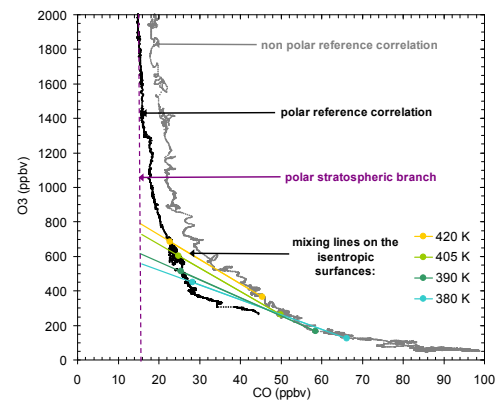


Figure 7. O_3 -CO correlation derived from SPIRALE measurements on 7 June 2005 at Teresina (grey line) and on 24 August 2009 at Kiruna (black line), the vertical dashed line represents the polar stratospheric branch and the color lines are examples of mixing lines at 420K, 405K, 390K and 380K.

CO is mainly produced or emitted in the troposphere and smoothly decreases with altitude. Unlike CO, O_3 is produced in the stratosphere and in low concentration in the troposphere. Consequently, without mixing between the stratosphere and the troposphere, the correlation between CO and O_3 looks like an “L”. The vertical

line is called stratospheric branch and the horizontal branch is called tropospheric branch. According to Fisher et al. [25], the mixing between tropospheric mid-latitude air and stratospheric air is represented by mixing lines between the tropospheric and the stratospheric branch. In our case, the transport between the non polar troposphere and the polar stratosphere occurs via the isentropic surfaces and the mixing lines connect the polar stratospheric branch, the polar correlation and the mid-latitude reference correlation at the same isentropic surface level (θ). Fig. 7 shows an example of the mixing line of the isentropic surfaces at 420 K, 405 K, 390 K and 380 K. Then, we can calculate the percentage of tropospheric mid-latitude air that is present in the polar stratosphere for both flights of SPIRALE. For each mixing line at an isentropic surface level (θ) fixed, we made the ratio between the distance polar stratospheric branch polar correlation and the distance polar stratospheric branch and non polar reference correlation. The results are shown in the Fig. 8 as a function of θ and compared with the vertical profile of CO for both flights. The shapes of the profiles are consistent with the proportion of non polar air: more than 40% between 380 K (~ 14 km) and 345 K (~ 12 km) for the 7 August 2009 flight compared to $\sim 20\%$ on 24-25 August 2009.

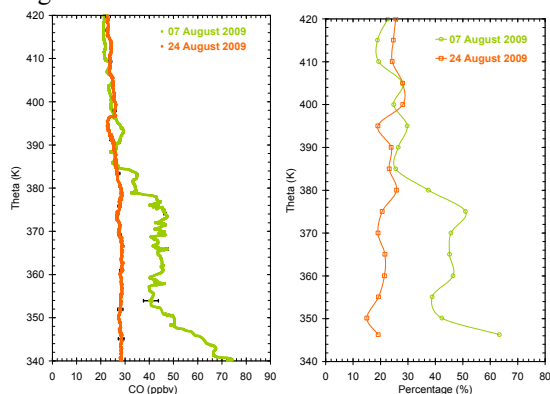


Figure 8. Right: Percentage of recent air in the mixed parcels sampled by SPIRALE on 7 August and on 24 August 2009 over Kiruna (Sweden) as a function of potential temperature, compared to the corresponding vertical profiles (Left).

4. CONCLUSION AND PERSPECTIVES

SPIRALE balloon-borne instrument observations over North Sweden during the summer 2009 showed a layer of air with very high carbon monoxide concentrations (mixing ratio near 80 ppbv) that is directly coupled with poleward intrusions from the mid-latitude on 7 August 2009. The layer extends over the potential temperature range 380 – 328 K (14 km – 10 km), with the maximum of CO at 330 K (11.5 km). These CO data are consistent with the MIMOSA model showing a region of low PV in the range 340-380 K. The CO

mixing ratio in the range 50 – 80 ppbv is typical of mid-latitude concentration. Backward trajectories at 11130 m and 11630 m from Kiruna revealed that the air had originated 10 days previously from North-East China and North-East Russia. The transport was also analyzed and revealed the important role of the long scale trans-Pacific transport and in particular the role of the jet stream. In this study, only one case was described. However, during the August 2009, there were three other similar phenomena simulated by MIMOSA model, on 21 August, on 26 August and on 28 August 2009. Therefore, the intrusion on 7 August 2009 seems not to be a particular case. The climatology of the intrusions during several summers should be established as well as the impact of Asia pollution on the Arctic regions. Studies have already started, led by several authors, such as Harrigan et al. [4]. They studied the impact of 3 regions of the world, Europe, USA and East China to Arctic by using trajectories model.

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